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THE DIABATIC WIND AND
TEMPERATURE PROFILES

by

H. Chuang and J. E. Cermak

Prepared for

U. S. Army Research Grant

DA-AMC-28-043-65-G20



ENGINEERING RESEARCH

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Fluid Dynamics and Diffusion Laboratory
College of Engineering
Colorado State University
Fort Collins, Colorado

December 1969

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ABSTRACT

KEYPS model and Monin-Obukhov's log-linear model were examined pertaining to their adequacy of describing wind and temperature profiles in thermally stratified shear flows for diversified thermal stability. The dimensionless wind shear and lapse rate for all ranges of thermal stability studied, $-2.0 \leq Ri \leq 0.4$, were shown to be linearly dependent on the dimensionless height derived from the log-linear model. Deacon numbers behaved quite differently from what were predicted by KEYPS model.

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INTRODUCTION

Wind and temperature profiles in thermally stratified shear flows are of great interest not only to the atmospheric scientist but also to the fluid dynamicist. Monin-Obukhov's well-known similarity theory (1) leads to the log-plus-linear profiles for flows in near-neutral conditions. Chuang and Cermak (2) used the field and laboratory data to show that wind and temperature profiles obtained in the laboratory as well as in the field approximately conform to the log-plus-linear model. A limited improvement of data scatter was made by adding a quadratic term to the log-plus-linear model (3). Bernstein (4) examined the existing three wind profile hypotheses--the log-linear profile, the KEYPS profile (5), and the exponential profile (6), and concluded that presently available data measured at O'Neill, Nebraska (7) and Kerang, Australia (8) were not sufficient to verify or refute any one of the three hypotheses. Pandolfo (9) introduced a free-convection model and claimed that it described the observed wind profiles quite accurately. However, he indicated that the free-convection profile was indistinguishable from the KEYPS profile.

Since KEYPS model is for interpolating the free and forced convection profiles, it is imperative to examine if this model can be used to describe wind and temperature profiles in all ranges of thermal stability. Monin-Obukhov's log-linear model has been proven (2) to be appropriate, even though not the most accurate method, for describing wind and

temperature profiles in thermally stratified shear flows of diversified stability. Therefore, these two models are examined in this paper. As to the free-convection model, which is nothing more than one of the "power law" profiles, it will be examined and presented in a separate paper.

Deacon numbers are defined and examined for the above-mentioned two models. The rate of change of eddy Prandtl number for different thermal stabilities is also studied.

BASIC EQUATIONS

The dimensionless wind shear, S , and lapse rate, R , are defined (5) as

$$S = \frac{kZ}{u_*} \frac{\partial U}{\partial Z} ,$$

and

$$R = \frac{\alpha Z}{T_*} \frac{\partial T}{\partial Z}$$

where k is von Karman's constant; u_* , the friction velocity; U , mean wind velocity; Z , height; T_* , the friction temperature; T , mean air temperature, and α , the ratio of eddy conductivity, K_h , to eddy viscosity, K_m , defined in the following form

$$\alpha = \frac{K_h}{K_m} = \frac{\overline{tw}}{\overline{uw}} \frac{\partial U}{\partial Z} / \frac{\partial T}{\partial Z} = 1/\text{eddy Prandtl number}$$

R becomes identical with S if the above equation inserted for α and if $u_*^2 = -\overline{uw}$ and $u_* T_* = -\overline{tw}$. In Reynolds analogy α is equal to unity.

When the mean wind velocity and temperature profiles are measured, the magnitudes of u_*/k and T_*/α can be estimated by means of the regression theory. Consequently, the dimensionless wind shear and lapse rate can be calculated to a moderate degree of accuracy provided that the velocity and temperature gradients in the vertical direction are accurately measured. This is not difficult to achieve in the laboratory where a continuous profile can be obtained

easily on an x-y plotter but is always a problem to face in the field. However, the method of interpolation can be used to find the velocity or temperature between two measured points and a more accurate gradient can then be determined.

KEYPS equation is an interpolation equation between free and forced convection. This equation was established in order to describe the gradual transition from forced to free convection. Hence, it cannot be expected to hold true in all ranges of thermal stability. KEYPS equation can be written in terms of Richardson number, Ri , as follows:

$$S = (1 - \gamma' Ri)^{-1/4} ,$$

where γ' is an arbitrary constant which can be determined by means of the least squares method when S and Ri are known. Similarly, the dimensionless lapse rate as defined previously will be identical with the above equation. This implies that the dimensionless wind shear and lapse rate are only a function of Richardson number.

Deacon numbers are defined as

$$DEU = - \frac{d(\ln \frac{\partial U}{\partial Z})}{d(\ln Z)} = 1 - \frac{d(\ln S)}{d(\ln Z)} , \quad (1)$$

and

$$DET = - \frac{d(\ln |\frac{\partial T}{\partial Z}|)}{d(\ln Z)} = 1 - \frac{d(\ln R)}{d(\ln Z)} . \quad (2)$$

When the dimensionless wind shear and lapse rate profiles are known the above dimensionless numbers can be easily

calculated. However, both Deacon numbers depend largely on the functional form of the dimensionless wind shear and lapse rate. When KEYPS equation is used for the dimensionless wind shear and lapse rate, both Deacon numbers will assume the same form as follows:

$$DEU = DET = 1 - \frac{\gamma' Ri}{4 - 3\gamma' Ri} \quad , \quad (3)$$

where Deacon numbers are unity at $Ri = 0$ (neutral flows) and approach $4/3$ as Richardson number approaches negative infinity. Since γ' may be different for different velocity and temperature profiles, it is conceivable that γ' becomes a parameter in the plot of Deacon numbers versus Richardson number.

In the log-plus-linear model, the mean wind velocity and temperature are written as

$$U - U_0 = A_1 \ln Z + B_1 Z + C_1,$$

and

$$T - T_0 = A_2 \ln Z + B_2 Z + C_2,$$

where $A_1 = u_* k$, $A_2 = T_*/\alpha$, the B's are constants divided by a length scale, L' , and the C's are constants. The dimensionless wind shear and lapse rate, then, assume the following form:

$$S = 1 + \frac{B_1}{A_1} Z \quad , \quad (4)$$

and

$$R = 1 + \frac{B_2}{A_2} Z . \quad (5)$$

Since S and R are identical, according to the similarity hypothesis, B_1/A_1 should be equal to B_2/A_2 . Assuming that

$$B_1/A_1 = B_2/A_2 = \beta'/L' ,$$

where β' is an arbitrary constant, and using the identity that $SRi \equiv Z/L'$ the above equations can be rewritten as

$$S = R = (1 - \beta'Ri)^{-1} \quad (6)$$

Equation (6) is similar in form to KEYPS equation except for the difference in power. Combining Eqs. (1) and (4) yields $DEU = 1/S$. Similarly $DET = 1/R$. Therefore, Deacon numbers assume the following form:

$$DEU = DET = 1 - \beta'Ri \quad (7)$$

Equation (7) is valid in both stable and unstable thermal stratification since Eqs. (4) and (5) hold true in all stability. It shows that Deacon numbers are linearly dependent on Richardson number under log-plus-linear model.

The relative rate of change of α in α , where α has been defined previously, can be shown to be a function of Deacon numbers, DEU and DET , as follows:

$$d\alpha/\alpha = (DET - DEU) dZ/Z$$

If Eqs. (4) and (5) are indeed identical, then the above equation implies that the relative rate of change of α in α should vanish, namely α should be constant in the whole profile under consideration. Similarity between the mean wind velocity and temperature profiles is not always exact but is an approximation. Therefore, B_1/A_1 is only approximately equal to B_2/A_2 , and the above equation is rewritten as

$$\frac{d\alpha/\alpha}{dZ/Z} = \left(\frac{B_1}{A_1} - \frac{B_2}{A_2} \right) L' Ri \quad (8)$$

The relative rate of change of Richardson number can also be shown to be a function of Deacon number as follows:

$$dRi/Ri = (2DEU - DET) dZ/Z - dT/T$$

The above equation can be rewritten as

$$\frac{dRi/Ri}{dZ/Z} = 2 DEU - DET - \frac{dT/T}{dZ/Z} \quad (9)$$

The length scale is defined as

$$L' = \frac{T_m}{g} \frac{u_* (\partial U / \partial Z)}{k (\partial T / \partial Z)} \approx \frac{T_m}{g} \frac{A_1^2}{A_2},$$

where T_m is the mean temperature, in absolute temperature scale, averaged over the whole profile and g is the gravitational acceleration.

EXPERIMENTAL RESULTS AND DISCUSSION

Mean wind velocity and temperature profiles in a wind-tunnel boundary layer were measured and reported by Chuang and Cermak (3). The flow over a horizontal flat plate was made thermally stable or unstable by cooling the plate and heating the air stream (inversion) or vice versa (lapse). For the purpose of easy handling, data were taken at closely spaced equidistant points from the continuous profiles of mean wind velocity and temperature. The range of height considered here is from 0.5 up to 8.2 cm., about one-seventh the total boundary layer thickness, with data points 0.7 cm apart. Figure 1 shows the results of log-plus-linear profile by means of regression theory and least squares method. Data of Project Prairie Grass (7) were calculated by means of interpolation between heights from 25 up to 750 cm at an equidistant interval of 50 cm, except the first increment being 25 cm, and were fitted to the profile by the least squares method. These results were also shown in the same figure where it was defined that

$$RUU = k(U_i - U_m)/u_* - \beta'(Z_i - Z_m)/L' ,$$

$$RTT = \alpha(T_i - T_m)/T_* - \beta'(Z_i - Z_m)/L' ,$$

and

$$RZZ = \ln Z_i - \frac{1}{N} \sum_{i=1}^N \ln Z_i$$

The subscript i refers to data point at i where $1 \leq i \leq N$ with total data points N , and the subscript m refers to the ensemble average over the profile of the named variables.

The dimensionless wind shear, lapse rate and Richardson number can be computed numerically when the derivatives of mean wind velocity and temperature with respect to the height are approximated by the ratios of finite increments of mean velocity and temperature about a stepwise sequence of points to that of height. The data-point interval, which is 0.7 cm for the laboratory data and 50 cm for the field data, is small and there is no abrupt change in both velocity and temperature profiles so that the approximation is fairly good. According to KEYPS equation, the dimensionless wind shear and lapse rate are a function of Ri only and in the case of dimensionless wind shear the constant γ' is equal to 18 as reported by Panofsky, et al (10). Figure 2 shows both the laboratory and field data of these two dimensionless quantities versus Richardson number. For comparison, Eqs. (4) and (5) were plotted in Fig. 3, where A's and B's were determined from the velocity and temperature profiles by the regression theory described in the last paragraph. It is evident that the diabatic wind and temperature profiles are better represented by Eqs. (4) and (5) than by KEYPS equation. The constant, γ' for both laboratory and field data is not necessarily equal to 18 but varies with height, wind and temperature profiles, and stability even though the over-all average value of it for

the field data may be 18. However, the variance of γ' is quite high.

Deacon numbers were computed according to Eqs. (1) and (2) and the finite difference approximation. They were plotted against Richardson number as shown in Fig. 4. Consider only the unstable cases where $Ri < 0$. Equation (3) is hardly descriptive of either laboratory or field data. This is due not only to the fact that KEYPS equation can vary with γ' as a parameter but also that the Deacon numbers are the second order derivatives of the diabatic wind and temperature profiles with respect to height Z . As the order of derivatives of a mean profile with respect to height increases, accuracy of the derivatives will decrease accordingly. Therefore, Deacon numbers will definitely amplify the scattering of data. Figure 5 shows the relationship between Deacon numbers and $\beta'Ri$. It reveals that Deacon numbers may not be exactly linearly dependent on Richardson number as predicted by Eq. (7).

Figure 6 shows the relative rate of change of α in α . Since $d\alpha/\alpha$ was reduced from Deacon numbers its accuracy was bad. However, this may show the general picture of similarity between the mean wind velocity profile and the mean temperature profile. It is imperative to measure the eddy Prandtl number directly. Figure 7 shows the effect of relative temperature gradient on the relative rate of change of Richardson number. The relative rate of change of Richardson number was computed by using the finite difference of Richardson number in a finite increment of height from the boundary.

CONCLUSIONS

The dimensionless wind shear and lapse rate for all ranges of thermal stability do not necessarily follow the prediction made by KEYPS equation. However, they are linearly dependent on the dimensionless height. Deacon numbers behave quite differently from what can be predicted by KEYPS model.

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LIST OF SYMBOLS

<u>Symbol</u>	<u>Definition</u>
A	$A_1 = u_*/k$, cm/sec; $A_2 = T_*/\alpha$, °C
B	Arbitrary constants
c_p	Specific heat of air at constant pressure, calories/°C/gm.
C	Arbitrary constants
g	Gravitational acceleration, cm/sec ²
H	Heat flux in the vertical direction, calories/cm ² /sec
k	von Karman constant
k_h	Eddy thermal conductivity, cm ² /sec
k_m	Eddy viscosity, cm ² /sec
L	Monin-Obukhov length scale, $L' = \alpha L$, cm
N	Total number of data collected in a profile
R	Dimensionless lapse rate
Ri	Richardson number, $g \frac{\partial T}{\partial Z} / [T(\frac{\partial U}{\partial Z})^2]$
S	Dimensionless wind shear
T	Mean absolute temperature, °K
T_*	$-H/(c_p \rho k u_*)$, friction temperature, °C
u_*	Friction velocity, cm/sec
U	Local mean velocity, cm/sec
Z	Height, cm
α	K_h/K_m
β'	Arbitrary constant
γ'	Arbitrary constant
ρ	Density of air, gm/cm ³

LIST OF SYMBOLS (Continued)

<u>Symbol</u>	<u>Definition</u>
$()_i$	The variable at height z_i
$()_m$	Mean value averaged over the profile
$()_o$	The variable at height z_o , an equivalent roughness height

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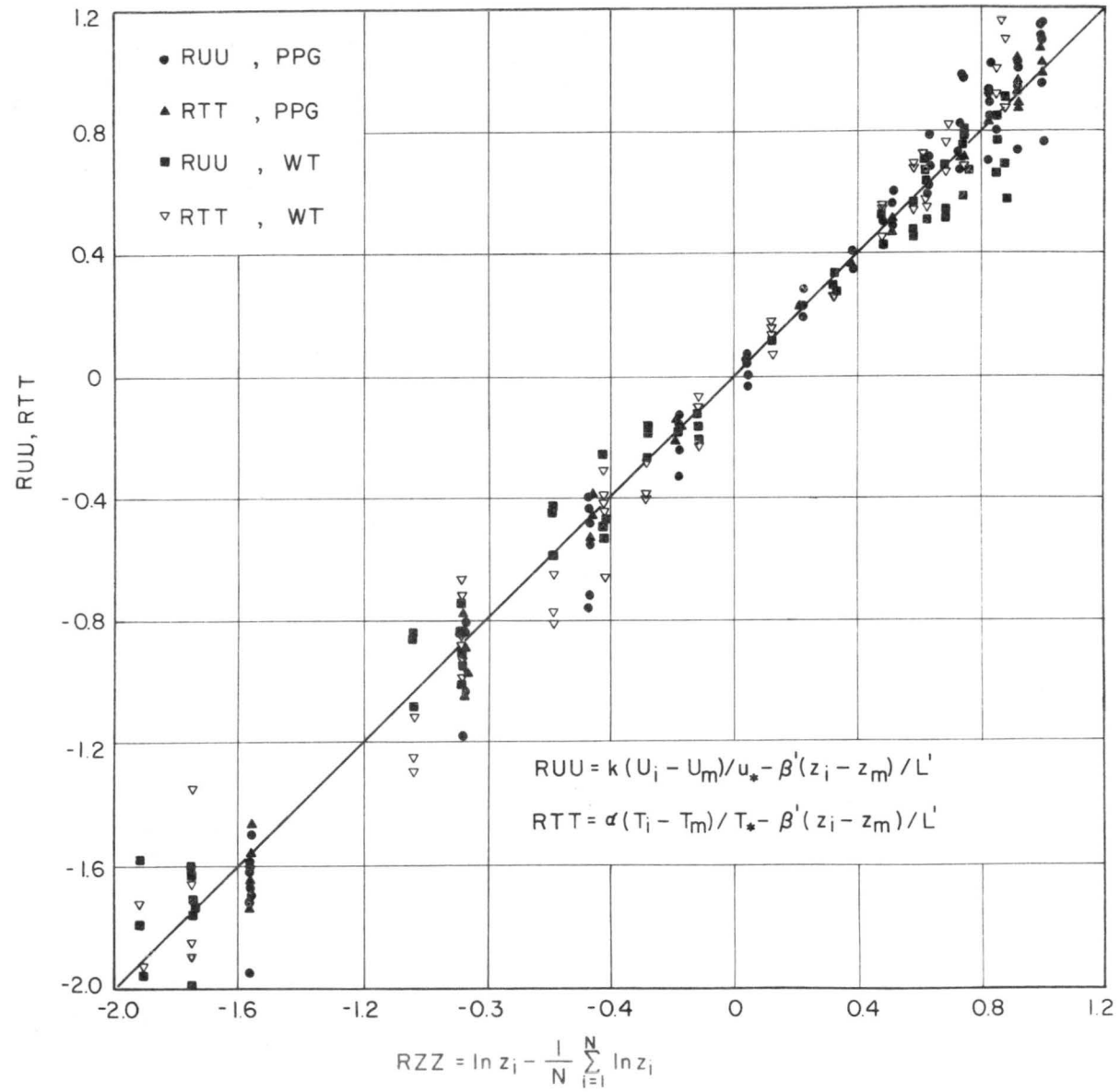


Fig. 1. Similarity law profile of mean velocity and temperature for heights from 0.5 to 8.2 cm (WT) or from 25 to 750 cm (PPG).

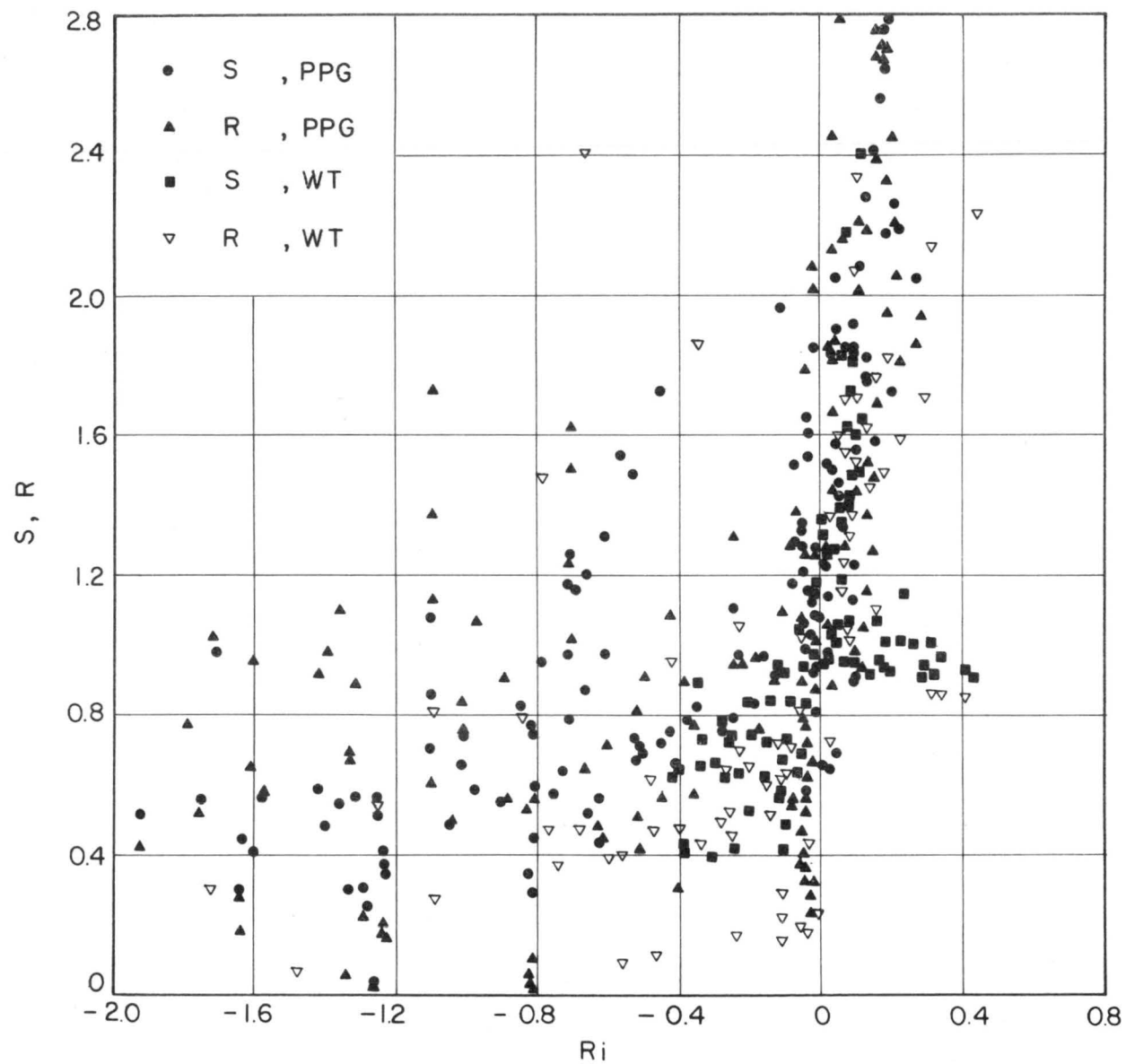


Fig. 2. Dimensionless wind shear and lapse rate versus Richardson number.

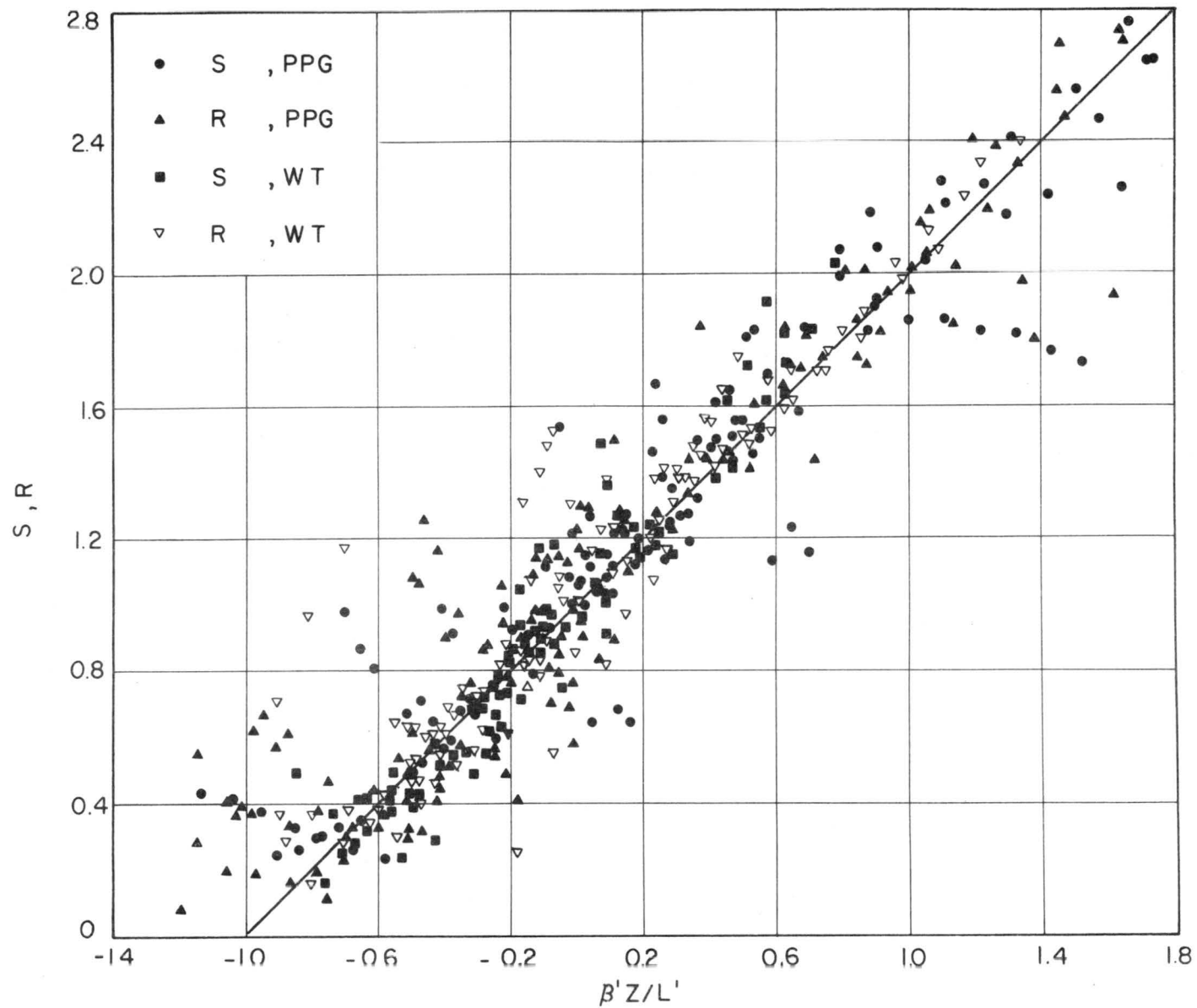


Fig. 3. Dimensionless wind shear and lapse rate versus dimensionless height $\beta'Z/L'$. The solid line represents Eqs. 4 and 5.

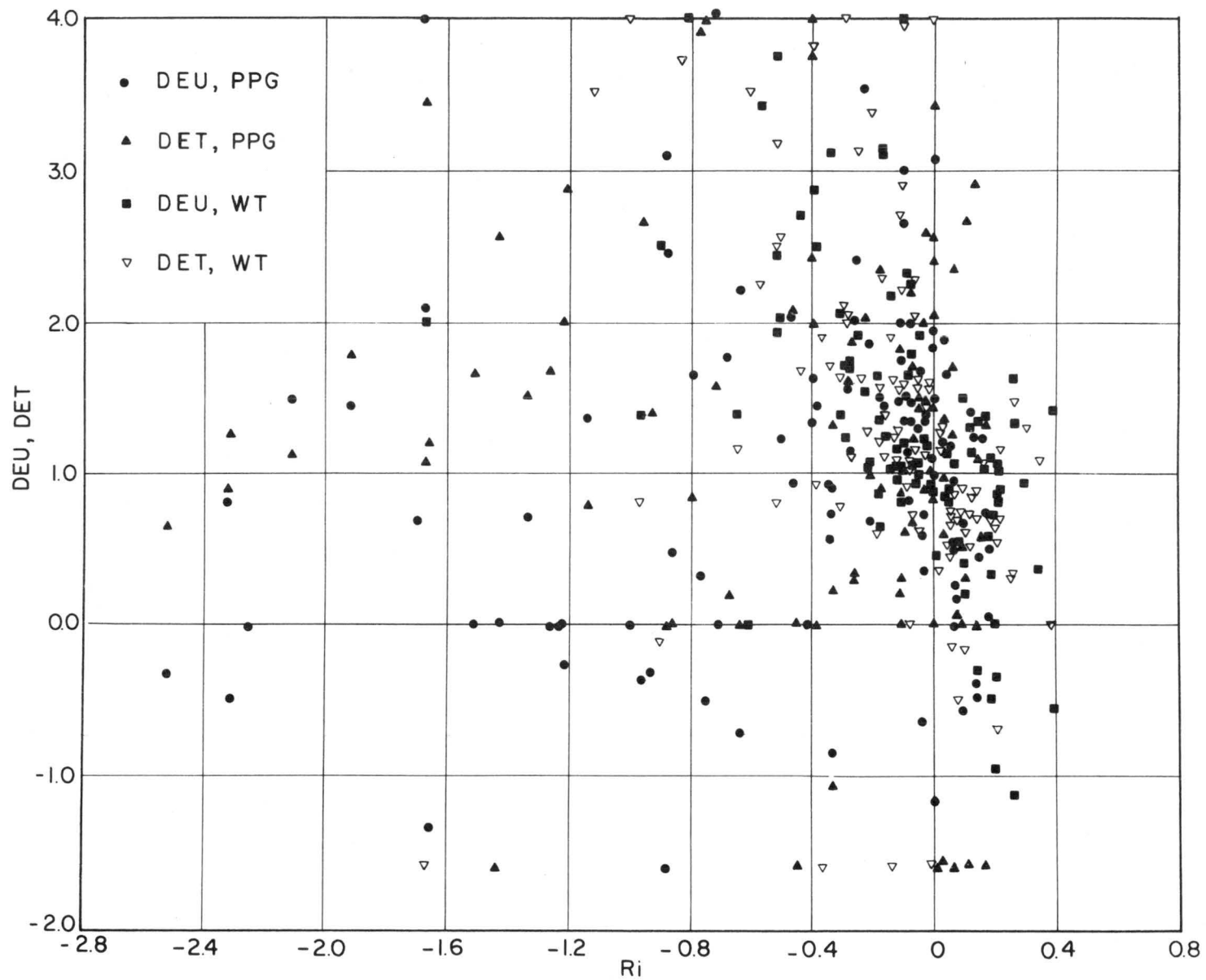


Fig. 4. Deacon numbers versus Richardson number.

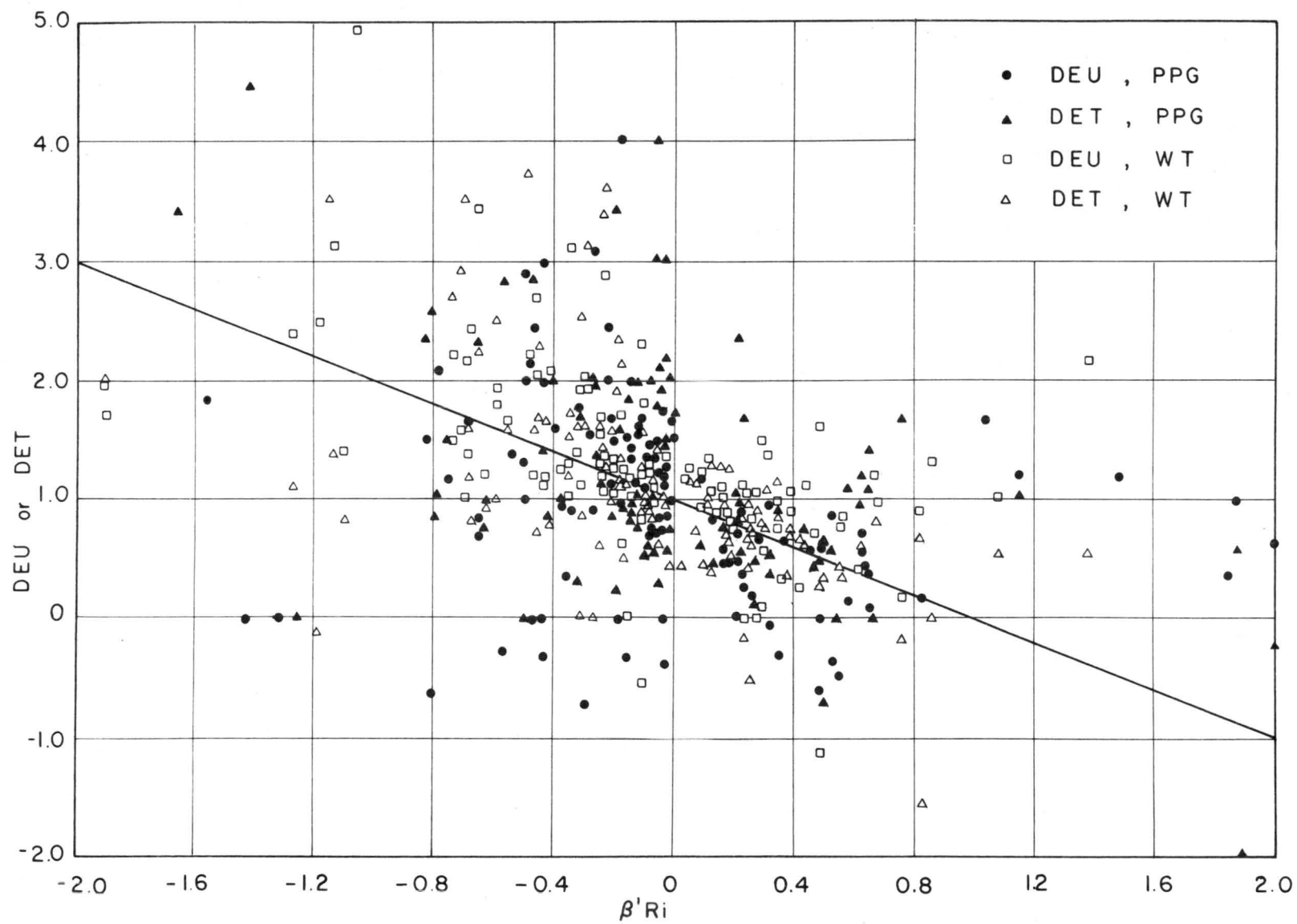


Fig. 5. Deacon numbers in terms of a linear function of Richardson number.
 The solid line represents Eq. 7.

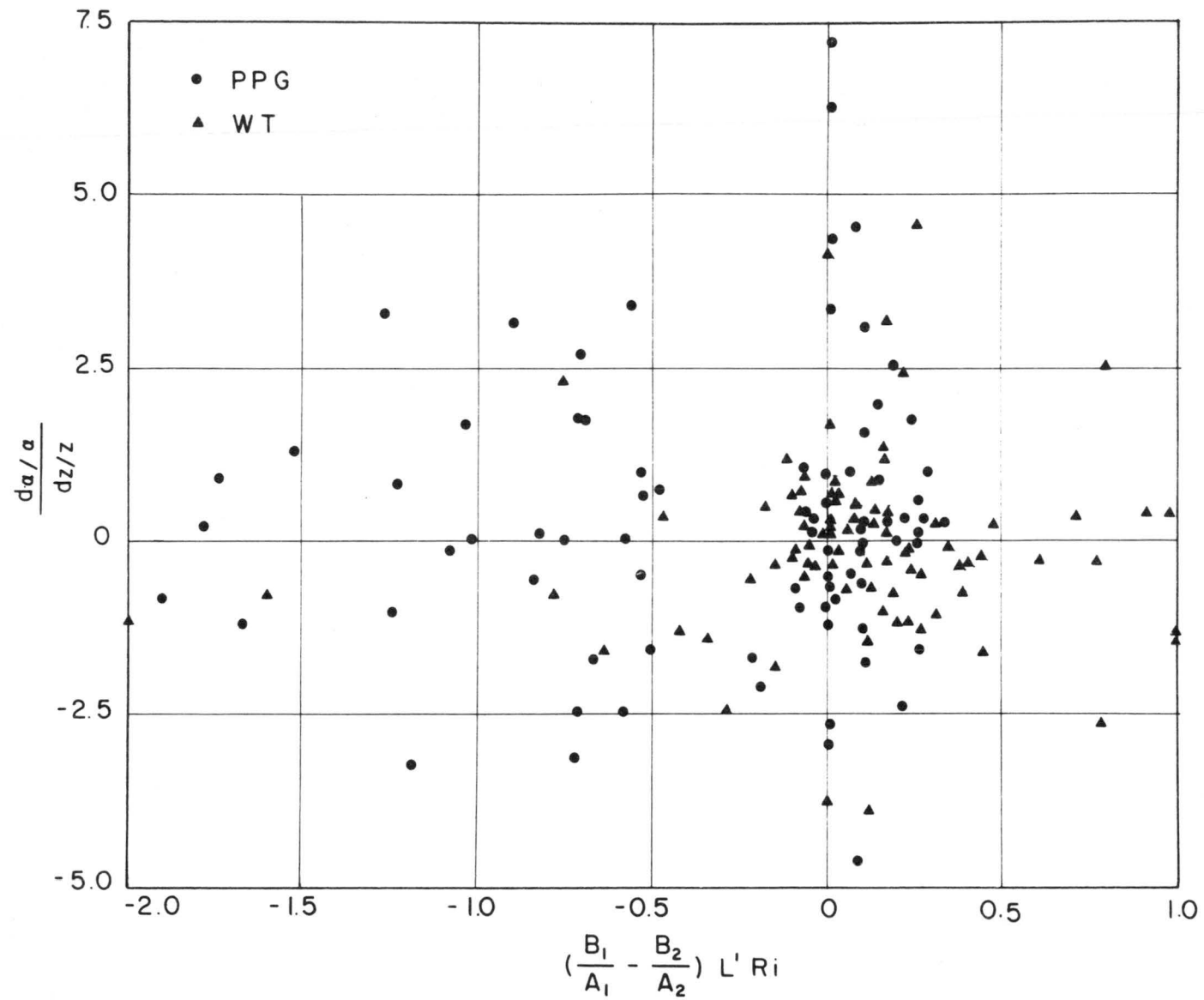


Fig. 6. Relative rate of change of α in terms of dimensionless parameters.

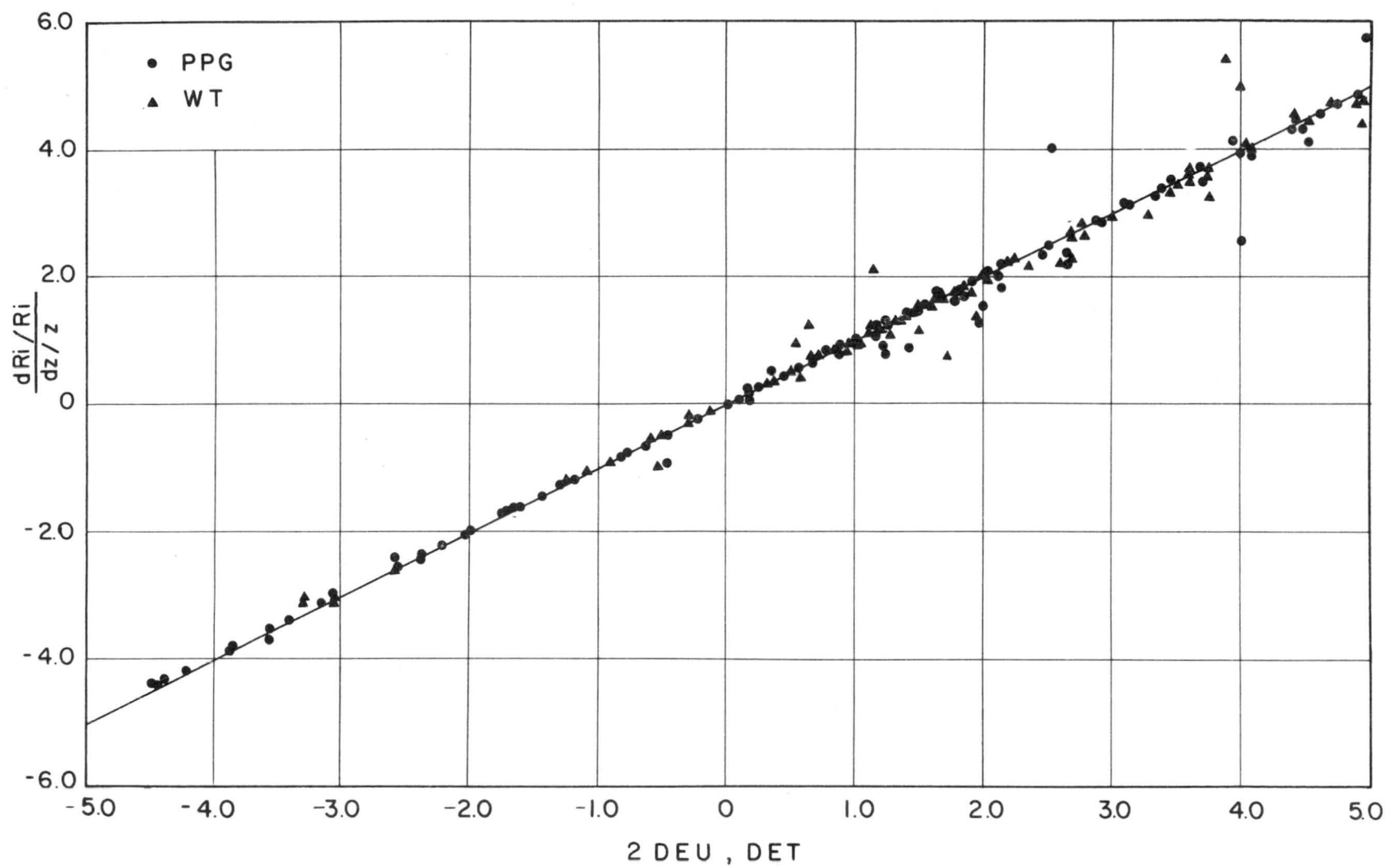


Fig. 7. Relative rate of change of Richardson number in terms of Deacon numbers.

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